"What is data assimilation really solving? What can it solve? How is the calculation actually done?"

Ichiro Fukumori (JPL)

GODAE Summer School September, 2004

What are we really solving? How can we solve it?

OUTLINE

- 1. What is assimilation really solving?
	- a) Filtering vs Smoothing
	- b) Observability & Controllability
	- c) Data Error and Model Error
- 2. Examples and Practical Issues; e.g., ECCO
	- a) A hierarchical approach
	- b) Practical derivation and implementation
	- c) Assessment

What is assimilation really solving?

Filtering vs Smoothing

- 1. Kalman filtering and recursive smoothing are both least-squares inversions, but solve different parts of the assimilation problem,
- 2. Filtered solutions are physically inconsistent,
- 3. Smoothed solutions can be physically consistent,
- 4. Estimating model error sources is fundamental.

The Mathematical Problem of Assimilation

and model errors)

 $\overline{\mathcal{L}}$

Assimilation problem: Given observations, y, and a model (A, G, H), what are the model states (x) and their controls (u)?

The Mathematical Problem of Assimilation

Least-Squares Solution

Almost all direct assimilation methods can be recognized as a form of least-squares.

For a linear inverse problem

$$
Ea \approx b
$$

The least-squares solution is given by (Gauss-Markov inversion)

$$
\hat{\mathbf{a}} = \mathbf{a}_0 + \mathbf{R}_{aa} \mathbf{E}^T \left(\mathbf{E} \mathbf{R}_{aa} \mathbf{E}^T + \mathbf{R}_{bb} \right)^{-1} \left(\mathbf{b} - \mathbf{E} \mathbf{a}_0 \right)
$$

$$
\boldsymbol{R}_{\hat{a}\hat{a}} = \boldsymbol{R}_{aa} - \boldsymbol{R}_{aa} \boldsymbol{E}^T \left(\boldsymbol{E} \boldsymbol{R}_{aa} \boldsymbol{E}^T + \boldsymbol{R}_{bb} \right)^{-1} \boldsymbol{E} \boldsymbol{R}_{aa}
$$

: least-squares solution ˆ**a**

 \mathbf{a}_0 , \mathbf{R}_{aa} : first guess (prior solution) and its error covariance matrix

R_{bb}: error covariance matrix of **b**

a_a: Error covariance of ˆ $\mathbf{R}_{\hat{\mathbf{a}}\hat{\mathbf{a}}}$: Error covariance of $\hat{\mathbf{a}}$

Kalman filter as Least-Squares

Linear inverse problem and its least-squares solution;

$$
\hat{\mathbf{a}} = \mathbf{a}_0 + \mathbf{R}_{aa} \mathbf{E}^T \left(\mathbf{E} \mathbf{R}_{aa} \mathbf{E}^T + \mathbf{R}_{bb} \right)^{-1} \left(\mathbf{b} - \mathbf{E} \mathbf{a}_0 \right)
$$
\n
$$
\mathbf{R}_{\hat{a}\hat{a}} = \mathbf{R}_{aa} - \mathbf{R}_{aa} \mathbf{E}^T \left(\mathbf{E} \mathbf{R}_{aa} \mathbf{E}^T + \mathbf{R}_{bb} \right)^{-1} \mathbf{E} \mathbf{R}_{aa}
$$
\nKalman filter estimate :
\n"forecast"
$$
\xrightarrow{\text{``forecast''}} \mathbf{\hat{x}}_t^f = \mathbf{A} \hat{\mathbf{x}}_{t-1}^a + \mathbf{G} \mathbf{u}_{t-1}
$$
\n"andysis"
$$
\xrightarrow{\text{``analysis''}} \mathbf{\hat{x}}_t^d = \hat{\mathbf{x}}_t^f + \mathbf{P}_t^f \mathbf{H}^T \left(\mathbf{H} \mathbf{P}_t^f \mathbf{H}^T + \mathbf{R}_t \right)^{-1} \left(\mathbf{y}_t - \mathbf{H} \hat{\mathbf{x}}_t^f \right)
$$
\n
$$
\mathbf{P}_t^f = \mathbf{A} \hat{\mathbf{P}}_{t-1}^a \mathbf{A}^T + \mathbf{G} \mathbf{Q} \mathbf{G}^T
$$
\n
$$
\hat{\mathbf{P}}_t^a = \mathbf{P}_t^f - \mathbf{P}_t^f \mathbf{H}^T \left(\mathbf{H} \mathbf{P}_t^f \mathbf{H}^T + \mathbf{R}_t \right)^{-1} \mathbf{H} \mathbf{P}_t^f
$$

There is a direct correspondence between the Kalman filtering algorithm and least-squares; i.e., Kalman filtering is nothing more than an inversion of the observation equation/operator . **H**

Inverse Solution

$$
\hat{\mathbf{a}} = \mathbf{a}_0 + \mathbf{R}_{aa} \mathbf{E}^T \left(\mathbf{E} \mathbf{R}_{aa} \mathbf{E}^T + \mathbf{R}_{bb} \right)^{-1} \left(\mathbf{b} - \mathbf{E} \mathbf{a}_0 \right)
$$

is the linear inverse solution of $\mathbf{E} \mathbf{a} \approx \mathbf{b}$ for any of following; 1. Minimum (expected error) variance: if $\min \left(diag \left\langle (\mathbf{a} - \overline{\mathbf{a}})(\mathbf{a} - \overline{\mathbf{a}})^T \right\rangle \right)$ $\mathbf{R_{aa}} = \left\langle (\mathbf{a}_0 - \overline{\mathbf{a}})(\mathbf{a}_0 - \overline{\mathbf{a}})^T \right\rangle$, $\mathbf{R_{bb}} = \left\langle (\mathbf{b} - \mathbf{E}\overline{\mathbf{a}})(\mathbf{b} - \mathbf{E}\overline{\mathbf{a}})^T \right\rangle$. True a

2. Least-squares: if $\min \left((\mathbf{a} - \mathbf{a}_0)^T \mathbf{W_a} (\mathbf{a} - \mathbf{a}_0) + (\mathbf{b} - \mathbf{E} \mathbf{a})^T \mathbf{W_b} (\mathbf{b} - \mathbf{E} \mathbf{a}) \right)$ $\frac{1}{2}$, $\bf{W}_{h} = \bf{R}_{hh}^{-1}$ $\mathbf{W}_{\mathbf{a}} = \mathbf{R}_{\mathbf{aa}}^{-1}, \ \ \mathbf{W}_{\mathbf{b}} = \mathbf{R}_{\mathbf{bb}}^{-1}$

3. Maximum likelihood: maxif Gaussian probability distribution $\wp(a) = \exp(- (a - a_0)^T R_{aa}^{-1} (a - a_0) - (b - Ea)^T R_{bb}^{-1} (b - Ea)$ $X\big[\wp(\mathbf{a})\big]$

Least-squares does not necessarily assume Gaussian statistics.

Kalman filter as Least-Squares

$$
\begin{bmatrix}\n\vdots \\
\mathbf{Hx}_{t} \\
\mathbf{Hx}_{t+1} \\
\vdots \\
\mathbf{x}_{t+1} - \mathbf{A}\mathbf{x}_{t} - \mathbf{G}\mathbf{u}_{t} \\
\vdots \\
\mathbf{Hx}_{t+2} - \mathbf{A}\mathbf{x}_{t+1} - \mathbf{G}\mathbf{u}_{t+1}\n\end{bmatrix}\n\approx\n\begin{bmatrix}\n\vdots \\
\mathbf{y}_{t} \\
\vdots \\
\mathbf{0} \\
\vdots \\
\mathbf{0} \\
\vdots \\
\mathbf{H}\mathbf{x}_{t-1} \\
\mathbf{H}\mathbf{x}_{t} \\
\mathbf{x}_{t} - \mathbf{A}\mathbf{x}_{0} - \mathbf{G}\mathbf{u}_{0} \\
\vdots \\
\mathbf{x}_{t+1} - \mathbf{A}\mathbf{x}_{t} - \mathbf{G}\mathbf{u}_{t}\n\end{bmatrix}\n\approx\n\begin{bmatrix}\n\vdots \\
\mathbf{y}_{t} \\
\mathbf{y}_{t} \\
\mathbf{y}_{t} \\
\vdots \\
\mathbf{0} \\
\mathbf{0} \\
\vdots \\
\mathbf{0}\n\end{bmatrix}
$$

$$
\hat{\mathbf{x}}_t^a = \hat{\mathbf{x}}_t^f + \mathbf{P}_t^f \mathbf{H}^T \left(\mathbf{H} \mathbf{P}_t^f \mathbf{H}^T + \mathbf{R}_t \right)^{-1} \left(\mathbf{y}_t - \mathbf{H} \hat{\mathbf{x}}_t^f \right)
$$

- •This is an inversion of the observation equation but **NOT** an inversion of the model equations.
- The estimate is the least-squares solution for the end state*,* but not at intermediate times.
- The virtue of Kalman filtering is that these inverted end states do not have to be recomputed over the entire period every time new observations are obtained.

Kalman filter as Least-Squares

The temporal evolution of filtered estimates are physically inconsistent (e.g., budgets do not close).

 $\mathbf{\hat{x}}_t^{\scriptscriptstyle J} = \mathbf{A} \mathbf{\hat{x}}_{t-1}^a + \mathbf{G} \mathbf{u}_{t-1}$ *f* t \leftarrow t \leftarrow t *a* $\mathbf{Kalman\ filter\ estimate:} \hat{\mathbf{x}}_{t}^{f} = \mathbf{A}\hat{\mathbf{x}}_{t-1}^{a} + \mathbf{Gu}_{t-1}$

$$
\hat{\mathbf{x}}_t^a = \hat{\mathbf{x}}_t^f + \mathbf{P}_t^f \mathbf{H}^T \left(\mathbf{H} \mathbf{P}_t^f \mathbf{H}^T + \mathbf{R}_t \right)^{-1} \left(\mathbf{y}_t - \mathbf{H} \hat{\mathbf{x}}_t^f \right)
$$

Combining the two steps give,

$$
\hat{\mathbf{x}}_t^a = \mathbf{A}\hat{\mathbf{x}}_{t-1}^a + \mathbf{G}\mathbf{u}_{t-1} + \mathbf{P}_t^f \mathbf{H}^T \left(\mathbf{H}\mathbf{P}_t^f \mathbf{H}^T + \mathbf{R}_t \right)^{-1} \left(\mathbf{y}_t - \mathbf{H}\hat{\mathbf{x}}_t^f \right)
$$

Not associated with any particular physics (e.g., advection, mixing, external forcing) embodied in **A** and **G**.

Consistency of Temporal Evolution

I.Fukumori/JPL

Example: Inconsistencies in Atmospheric Analyses

24% of mass change in NCEP's operational analysis is physically **unaccounted for**.

change over 6-hours assimilation increment

Smother as Least-Squares

Kalman filter estimate:

$$
\hat{\mathbf{x}}_t^a = \mathbf{A}\hat{\mathbf{x}}_{t-1}^a + \mathbf{G}\mathbf{u}_{t-1} + \mathbf{P}_t^f \mathbf{H}^T \left(\mathbf{H}\mathbf{P}_t^f \mathbf{H}^T + \mathbf{R}_t \right)^{-1} \left(\mathbf{y}_t - \mathbf{H}\hat{\mathbf{x}}_t^f \right)
$$

Smoothing problem: Given the updated estimate $\hat{\mathbf{x}}_t^a$ invert the **model** to correct the prior state x_{*t*−1} and control u_{*t−1*} that; $\hat{\textbf{X}}_t^a$

$$
\hat{\mathbf{x}}_t^a = \mathbf{A}\hat{\mathbf{x}}_{t-1}^s + \mathbf{G}\hat{\mathbf{u}}_{t-1}^s = \begin{pmatrix} \mathbf{A} & \mathbf{G} \end{pmatrix} \begin{pmatrix} \hat{\mathbf{x}}_{t-1}^s \\ \hat{\mathbf{u}}_{t-1}^s \end{pmatrix}
$$

A least-squares solution that exactly satisfies the inverse problem $\mathbf{E}\mathbf{a} = \mathbf{b}$ can be obtained by setting $\mathbf{R}_{bb} = 0$ in the canonical least-squares solution;

By substitution, we have;

$$
\begin{pmatrix} \hat{\mathbf{x}}_{t-1}^{s} \\ \hat{\mathbf{u}}_{t-1}^{s} \end{pmatrix} = \begin{pmatrix} \hat{\mathbf{x}}_{t-1}^{a} \\ \mathbf{u}_{t-1} \end{pmatrix} + \begin{pmatrix} \mathbf{P}_{t-1}^{a} \mathbf{A}^{T} \left(\mathbf{A} \mathbf{P}_{t-1}^{a} \mathbf{A}^{T} + \mathbf{G} \mathbf{Q}_{t-1} \mathbf{G}^{T} \right)^{-1} \\ \mathbf{Q}_{t-1} \mathbf{G}^{T} \left(\mathbf{G} \mathbf{Q}_{t-1} \mathbf{G}^{T} + \mathbf{A} \mathbf{P}_{t-1}^{a} \mathbf{A}^{T} \right)^{-1} \end{pmatrix} \left(\hat{\mathbf{x}}_{t}^{a} - \mathbf{A} \hat{\mathbf{x}}_{t-1}^{a} - \mathbf{G} \mathbf{u}_{t-1} \right)
$$

Smother as Least-Squares

More generally,

$$
\frac{\begin{pmatrix} \hat{\mathbf{x}}_{t}^{s} \\ \hat{\mathbf{u}}_{t}^{s} \end{pmatrix} = \begin{pmatrix} \hat{\mathbf{x}}_{t}^{a} \\ \mathbf{u}_{t} \end{pmatrix} + \begin{pmatrix} \mathbf{P}_{t}^{a} \mathbf{A}^{T} \left(\mathbf{A} \mathbf{P}_{t}^{a} \mathbf{A}^{T} + \mathbf{G} \mathbf{Q}_{t} \mathbf{G}^{T} \right)^{-1} \\ \mathbf{Q}_{t} \mathbf{G}^{T} \left(\mathbf{G} \mathbf{Q}_{t} \mathbf{G}^{T} + \mathbf{A} \mathbf{P}_{t}^{a} \mathbf{A}^{T} \right)^{-1} \mathbf{A} \mathbf{P}_{t}^{a} \\ \mathbf{Q}_{t}^{s} \end{pmatrix}} = \begin{pmatrix} \mathbf{P}_{t}^{a} \\ \mathbf{Q}_{t} \end{pmatrix} - \begin{pmatrix} \mathbf{P}_{t}^{a} \mathbf{A}^{T} \left(\mathbf{A} \mathbf{P}_{t}^{a} \mathbf{A}^{T} + \mathbf{G} \mathbf{Q}_{t} \mathbf{G}^{T} \right)^{-1} \mathbf{A} \mathbf{P}_{t}^{a} \\ \mathbf{Q}_{t} \mathbf{G}^{T} \left(\mathbf{G} \mathbf{Q}_{t} \mathbf{G}^{T} + \mathbf{A} \mathbf{P}_{t}^{a} \mathbf{A}^{T} \right)^{-1} \mathbf{G} \mathbf{Q}_{t} \end{pmatrix} + \begin{pmatrix} \mathbf{L}_{t}^{a} \mathbf{P}_{t+1}^{s} \mathbf{L}_{t}^{a} \\ \mathbf{M}_{t}^{a} \mathbf{P}_{t+1}^{s} \mathbf{M}_{t}^{a} \\ \mathbf{M}_{t}^{a} \mathbf{P}_{t}^{s} \mathbf{M}_{t}^{a} \end{pmatrix}}{\mathbf{W}_{t}^{a} \equiv \mathbf{Q}_{t} \mathbf{G}^{T} \left(\mathbf{G} \mathbf{Q}_{t} \mathbf{G}^{T} + \mathbf{A} \mathbf{P}_{t}^{a} \mathbf{A}^{T} \right)^{-1} \qquad \text{Extra term due to errors in } \hat{\mathbf{x}}_{t+1}^{s} \\ \text{Least-squares solution;} \qquad \hat{\
$$

Smoother as Least-Squares

$$
\begin{pmatrix} \hat{\mathbf{x}}_t^s \\ \hat{\mathbf{u}}_t^s \end{pmatrix} = \begin{pmatrix} \hat{\mathbf{x}}_t^a \\ \mathbf{u}_t \end{pmatrix} + \begin{pmatrix} \mathbf{P}_t^a \mathbf{A}^T \left(\mathbf{A} \mathbf{P}_t^a \mathbf{A}^T + \mathbf{G} \mathbf{Q}_t \mathbf{G}^T \right)^{-1} \\ \mathbf{Q}_t \mathbf{G}^T \left(\mathbf{G} \mathbf{Q}_t \mathbf{G}^T + \mathbf{A} \mathbf{P}_t^a \mathbf{A}^T \right)^{-1} \end{pmatrix} \left(\hat{\mathbf{x}}_{t+1}^s - \mathbf{A} \hat{\mathbf{x}}_t^a - \mathbf{G} \mathbf{u}_t \right)
$$

• This is the Rauch-Tung-Striebel (RTS) fixed-interval smoother and is an inversion of the model equations.

$$
\begin{bmatrix}\n&\mathbf{Hx}_{t} \\
&\mathbf{Hx}_{t+1} \\
&\mathbf{Hx}_{t+1} \\
&\mathbf{Hx}_{t+1}\n\end{bmatrix}\n\approx\n\begin{bmatrix}\n\vdots \\
\mathbf{y}_{t} \\
\mathbf{y}_{t+1} \\
\vdots \\
\mathbf{y}_{t+1} \\
\mathbf{y}_{t+1} \\
\vdots\n\end{bmatrix}
$$

• This solves the entire assimilation problem by sequentially solving smaller calculations.

• Another virtue is its results' physical consistency.

Smoother as Least-Squares

$$
\begin{pmatrix} \hat{\mathbf{x}}_t^s \\ \hat{\mathbf{u}}_t^s \end{pmatrix} = \begin{pmatrix} \hat{\mathbf{x}}_t^a \\ \mathbf{u}_t \end{pmatrix} + \begin{pmatrix} \mathbf{P}_t^a \mathbf{A}^{\mathrm{T}} \left(\mathbf{A} \mathbf{P}_t^a \mathbf{A}^{\mathrm{T}} + \mathbf{G} \mathbf{Q}_t \mathbf{G}^{\mathrm{T}} \right)^{-1} \\ \mathbf{Q}_t \mathbf{G}^{\mathrm{T}} \left(\mathbf{G} \mathbf{Q}_t \mathbf{G}^{\mathrm{T}} + \mathbf{A} \mathbf{P}_t^a \mathbf{A}^{\mathrm{T}} \right)^{-1} \end{pmatrix} \begin{pmatrix} \hat{\mathbf{x}}_{t+1}^s - \mathbf{A} \hat{\mathbf{x}}_t^a - \mathbf{G} \mathbf{u}_t \end{pmatrix}
$$

 $\mathbf{\hat{x}}^s = \mathbf{A}\mathbf{\hat{x}}$

 $\hat{\mathbf{x}}_t^s = \mathbf{A}\hat{\mathbf{x}}_{t-1}^s + \mathbf{G}\hat{\mathbf{u}}_{t-1}^s$

 $1 \quad 1 \quad u_{t-1}$

ˆ

• By construction, this satisfies the model equation,

making it amenable to process studies. This should **not** be confused with the "strong constraint" solution; $\mathbf{In} \ \mathbf{fact}, \quad \hat{\mathbf{x}}^{\scriptscriptstyle S}_{\scriptscriptstyle t} \neq \mathbf{A}\hat{\mathbf{x}}^{\scriptscriptstyle S}_{\scriptscriptstyle t-1} + \mathbf{Gu}_{\scriptscriptstyle t-1}$

- Control estimates are of the essence. Smoothed state estimates can be computed from the control estimates.
- The model for model process noise \mathbf{u}_t must be physically based for the result to be physically sensible.

Smoother as Least-Squares

$$
\begin{pmatrix} \hat{\mathbf{x}}_t^s \\ \hat{\mathbf{u}}_t^s \end{pmatrix} = \begin{pmatrix} \hat{\mathbf{x}}_t^a \\ \mathbf{u}_t \end{pmatrix} + \begin{pmatrix} \mathbf{P}_t^a \mathbf{A}^{\mathrm{T}} \left(\mathbf{A} \mathbf{P}_t^a \mathbf{A}^{\mathrm{T}} + \mathbf{G} \mathbf{Q}_t \mathbf{G}^{\mathrm{T}} \right)^{-1} \\ \mathbf{Q}_t \mathbf{G}^{\mathrm{T}} \left(\mathbf{G} \mathbf{Q}_t \mathbf{G}^{\mathrm{T}} + \mathbf{A} \mathbf{P}_t^a \mathbf{A}^{\mathrm{T}} \right)^{-1} \end{pmatrix} \left(\hat{\mathbf{x}}_{t+1}^s - \mathbf{A} \hat{\mathbf{x}}_t^a - \mathbf{G} \mathbf{u}_t \right)
$$

- The RTS smoother recursively carries information of formally future observations backwards in time.
- Utilizes all data within a fixed time-interval whereas the Kalman filter only utilizes data up to that instant.

• More accurate than the filter's estimate.

Other Smoothers

Besides the recursive RTS smoother, the adjoint method and Green's function optimizations are other commonly used fixed-interval smoothers.

What are we really solving? How can we solve it?

OUTLINE

- 1. What is assimilation really solving?
	- a) Filtering vs Smoothing
	- b) Observability & Controllability
	- c) Data Error and Model Error
- 2. Examples and Practical Issues; e.g., ECCO
	- a) A hierarchical approach
	- b) Practical derivation and implementation
	- c) Assessment

Observability & Controllability

Kalman Filter is an inversion of observation **H**. RTS smoother is an inversion of model evolution **A** and **G**.

Identifies the problem the filter & smoother are solving (inverting).

Observability identifies if the model state that can be determined. Controllability identifies if controls can be determined.

Identifies the solution that could be determined.

Observability

If there are no errors in the observations and no errors in the model (except state is unknown), can we determine the model state uniquely ?

$$
\mathbf{y}_{1} = \mathbf{H}_{1}\mathbf{x}_{1} = \mathbf{H}_{1} \left(\mathbf{A}\mathbf{x}_{0} - \mathbf{G}\mathbf{u}_{0}\right) \longrightarrow \mathbf{y}_{1} + \mathbf{H}_{1}\mathbf{G}\mathbf{u}_{0} = \mathbf{H}_{2}\mathbf{A}\mathbf{x}_{0}
$$
\n
$$
\vdots
$$
\n
$$
\mathbf{y}_{t} + \mathbf{H}_{t}\mathbf{G}\mathbf{u}_{t-1} + \mathbf{H}_{t}\mathbf{A}\mathbf{G}\mathbf{u}_{t-2} + \cdots + \mathbf{H}_{t}\mathbf{A}^{t-1}\mathbf{G}\mathbf{u}_{0} \equiv \hat{\mathbf{y}}_{t} = \mathbf{H}_{t}\mathbf{A}^{t}\mathbf{x}_{0}
$$
\n
$$
\begin{pmatrix} \hat{\mathbf{y}}_{1} \\ \vdots \\ \hat{\mathbf{y}}_{t} \end{pmatrix} = \mathbf{O}\mathbf{x}_{1} \text{ where } \mathbf{O} \equiv \begin{pmatrix} \mathbf{H}_{1}\mathbf{A} \\ \vdots \\ \mathbf{H}_{t}\mathbf{A}^{t} \end{pmatrix}
$$

If **O** is full rank (column), there is a unique solution for the state, and the system is said to be observable.

Controllability

Can we determine controls that can drive the model to an arbitrary state?

$$
\mathbf{x}_{1} = \mathbf{A}\mathbf{x}_{0} + \mathbf{G}\mathbf{u}_{0}
$$
\n
$$
\mathbf{x}_{2} = \mathbf{A}\mathbf{x}_{1} + \mathbf{G}\mathbf{u}_{1} = \mathbf{A}^{2}\mathbf{x}_{0} + \mathbf{A}\mathbf{G}\mathbf{u}_{0} + \mathbf{G}\mathbf{u}_{1}
$$
\n
$$
\vdots
$$
\n
$$
\mathbf{x}_{t} = \mathbf{A}\mathbf{x}_{t-1} + \mathbf{G}\mathbf{u}_{t-1} = \mathbf{A}^{t}\mathbf{x}_{0} + \mathbf{A}^{t-1}\mathbf{G}\mathbf{u}_{0} + \mathbf{A}^{t-2}\mathbf{G}\mathbf{u}_{1} + \mathbf{G}\mathbf{u}_{t-1}
$$
\n
$$
\mathbf{x}_{t} - \mathbf{A}^{t}\mathbf{x}_{0} = \mathbf{C} \begin{pmatrix} \mathbf{u}_{0} \\ \mathbf{u}_{1} \\ \vdots \\ \mathbf{u}_{t-1} \end{pmatrix} \text{ where } \mathbf{C} \equiv (\mathbf{G} \quad \mathbf{A}\mathbf{G} \quad \cdots \quad \mathbf{A}^{t-1}\mathbf{G})
$$

If **C** is full rank (row), there is a solution for the control that satisfies this equation, and the system is said to be controllable.

Observability & Controllability

These conditions are conceptually important. But the physical state being observable does not mean any meaningful estimate can be made; i.e., Difference between P^a and P^f may be indistinguishable.

What are we really solving? How can we solve it?

OUTLINE

- 1. What is assimilation really solving?
	- a) Filtering vs Smoothing
	- b) Observability & Controllability
	- c) Data Error and Model Error
- 2. Examples and Practical Issues; e.g., ECCO
	- a) A hierarchical approach
	- b) Practical derivation and implementation
	- c) Assessment

What is assimilation really solving?

"Data Error" and "Model Error"

- 1. The two are better thought of as observation constraint error and model constraint error, respectively, as opposed to errors of the data and model, per se.
- 2. "Data Error" includes representation errors, i.e., things that the models do not represent,
- 3. How to specify the two prior errors.

"Data Error" & "Model Error"

"Data Error"

R represents the error of the data constraint and includes model representation error.

The true model state $\overline{\mathbf{x}}_t$ is a finite dimensional representation of the true state of the ocean, \mathbf{w}_t

 $\overline{\mathbf{x}}_t$ ≡ $\mathbf{\Pi}\mathbf{w}_t$ $\qquad \mathbf{\Pi}$: Function defining the model state.

Observations are related to the true state of the ocean by

 $y_t = E w_t + \varepsilon$

E: Function defining observation. ^ε: Measurement error

In terms of the true model state, (after Cohn, 1997, *J. Met. Soc. Jap.*)

$$
\boxed{\mathbf{y}_t = \mathbf{H}\overline{\mathbf{x}}_t + \left\{\mathbf{E}\mathbf{w}_t - \mathbf{H}\mathbf{\Pi}\mathbf{w}_t\right\} + \varepsilon}
$$

Error of observation equation $\mathbf{y}_{_t} = \mathbf{H}\mathbf{x}_{_t}$

Ew*^t ^t* [−] **^HΠ^w** : Representation error

"Data Error"

Ew*^t ^t* [−] **^HΠ^w** : Representation error

Error-free data Error-free model equivalent of data **Π** : Defines the model state.

- **E**: Defines observation.
- **H**: Defines model equivalent of data
- Examples;
- Meso-scale variability for a non-eddy resolving model,
- External tides for a rigid lid model,
- Skin temperature for model with thick surface layer,
- Micro-structure for most models,
- Barotropic variability for a reduced-gravity model,
- Baroclinic variability for a barotropic model,
- Meteorologist do not force models too close to observations in numerical weather forecasting,
- Individual float & drifter trajectories for any finite difference model.

Examples of Drifter Trajectory

(Paduan and Niiler, 1993, JPO)

Example of diverging float Typical "spaghetti" diagram
trajectories in the trajectories

"Model Error"

Q represents the error of the model constraint.

 $\mathbf{x}_{t+1} = A\left(\mathbf{x}_t, \hat{\mathbf{u}}_t^{\alpha}\right)$ $= A(\mathbf{x}, \hat{\mathbf{u}})$ $\overline{\mathbf{X}}_{t+1} = \mathbf{\Pi} \mathbf{W}_{t+1} = \mathbf{\Pi} L \big(\mathbf{W}_{t}, \mathbf{V}_{t} \big)$ $\mathbf{A}_{t} = A\left(\overline{\mathbf{X}}_{t}, \mathbf{\hat{u}}_{t}^{0}\right) + \left\{\mathbf{\Pi}L\left(\mathbf{W}_{t}, \mathbf{V}_{t}\right) - A\left(\mathbf{\Pi}\mathbf{W}_{t}, \mathbf{\hat{u}}_{t}^{0}\right)\right\}$ $\mathbf{H}_{t+1} = A(\overline{\mathbf{X}}_t, \mathbf{\hat{u}}_t^{\mathrm{o}}) + \{ \mathbf{\Pi} L(\mathbf{W}_t, \mathbf{V}_t) - A(\mathbf{\Pi}\mathbf{W}_t, \mathbf{\hat{u}}_t^{\mathrm{o}})\}$ $\overline{\mathbf{x}}_{t+1} = A(\overline{\mathbf{x}}_t, \hat{\mathbf{u}}_t^0) + \{ \Pi L(\mathbf{w}_t, \mathbf{v}_t) - A(\Pi \mathbf{w}_t, \hat{\mathbf{u}}_t^0) \}$ Ocean evolution: $\left(\mathbf{w}_{_{t}}, \mathbf{v}_{_{t}}\right)$ – $A\!\left(\mathbf{\Pi} \mathbf{w}_{_{t}}, \hat{\mathbf{u}}_{_{t}}^{_{0}}\right)$: Model Process Noise Model evolution: $\mathbf{x}_{_{t+1}} = A(\mathbf{x}_{_t}, \hat{\mathbf{u}})$ $\mathbf{w}_{t+1} \neq L(\mathbf{w}_{t}, \mathbf{v}_{t})$ The true model evolution can be written as, Forcing, boundary condition, mixing parameters, model error sources. Error of the model evolution; $\mathbf{x}_{_{t+1}} = A\!\left(\mathbf{x}_{_{t}}, \hat{\mathbf{u}}_{_{t}}^{\text{o}}\right)$ Functions describing the evolution $= A(\mathbf{x}_{t}, \hat{\mathbf{u}})_{t}^{\mathsf{U}}$ Real forcing & boundary condition. ˆ \mathbf{H} *L*(\mathbf{w}_t , \mathbf{v}_t) – A ($\mathbf{H}\mathbf{w}_t$, \mathbf{w}_t $-A($ **IIw**_t, $\hat{\mathbf{u}}_t^{\mathbf{c}}$ Model evolution given true model stateTrue evolution

"Model Error"

 $\left(\mathbf{w}_{_{t}}, \mathbf{v}_{_{t}}\right)$ – $A\!\left(\mathbf{\Pi w}_{_{t}}, \hat{\mathbf{u}}_{_{t}}^{_{0}}\right)$: Model Process Noise ˆ \mathbf{H} *L*(\mathbf{w}_t , \mathbf{v}_t) – A ($\mathbf{H}\mathbf{w}_t$, \mathbf{w}_t $-A($ **IIw**_t, $\hat{\mathbf{u}}_t^{\mathbf{c}}$

Examples;

- 1) Differences in $\hat{\textbf{u}}_t^0$ and \textbf{v} (errors in non-state variables); *forcing, boundary condition, model parameters,*
- 2) Errors due to differences in *A* and *L* (errors in model algorithm); *finite difference, interaction with scales and processes ignored*

$$
\overline{\mathbf{x}}_{t+1} = A(\overline{\mathbf{x}}_t, \hat{\mathbf{u}}_t^0) + \left\{ \Pi L(\mathbf{w}_t, \mathbf{v}_t) - A(\Pi \mathbf{w}_t, \hat{\mathbf{u}}_t^0) \right\}
$$

= $A(\overline{\mathbf{x}}_t, \overline{\mathbf{u}}_t)$

Collectively, sources of model process noise are modeled as elements of the model control vector.

Prior Error Specification

Covariance Matching (Fu et al., 1993, *JPO*)

Estimate **R** & **Q** from model-data comparison.

Data **y** and its model simulation equivalent **m** could be written as; $\mathbf{y} = \mathbf{s} + \mathbf{r}$

 $m = s + n$

where, s: true signal $(H\bar{x})$, r: data error, **m**: simulation of $\mathbf{y}\left(\mathbf{H}\mathbf{x}_{\textit{sim}}\right)$,

 $\mathbf{n}: \text{ simulation error } \left(\mathbf{H} \mathbf{x}_{\textit{sim}} - \mathbf{H} \mathbf{\overline{x}} \right)$

Then,

$$
\langle \mathbf{y}\mathbf{y}^T \rangle = \langle \mathbf{s}\mathbf{s}^T \rangle + \langle \mathbf{s}\mathbf{y}^T \rangle + \langle \mathbf{r}\mathbf{s}^T \rangle + \langle \mathbf{r}\mathbf{r}^T \rangle
$$

$$
\langle \mathbf{m}\mathbf{m}^T \rangle = \langle \mathbf{s}\mathbf{s}^T \rangle + \langle \mathbf{s}\mathbf{n}^T \rangle + \langle \mathbf{n}\mathbf{s}^T \rangle + \langle \mathbf{n}\mathbf{n}^T \rangle
$$

$$
\langle \mathbf{m}\mathbf{y}^T \rangle = \langle \mathbf{s}\mathbf{s}^T \rangle + \langle \mathbf{s}\mathbf{r}^T \rangle + \langle \mathbf{n}\mathbf{s}^T \rangle + \langle \mathbf{p}\mathbf{r}^T \rangle
$$

Prior Error Specification

If,
$$
\langle \mathbf{s} \mathbf{n}^T \rangle \approx 0
$$
 etc,
\n
$$
\mathbf{R} = \langle \mathbf{r} \mathbf{r}^T \rangle = \langle \mathbf{y} \mathbf{y}^T \rangle - \langle \mathbf{m} \mathbf{y}^T \rangle
$$
\n
$$
\mathbf{R} = \langle \mathbf{r} \mathbf{n} \rangle = \langle \mathbf{m} \mathbf{m}^T \rangle - \langle \mathbf{m} \mathbf{y}^T \rangle
$$
\n
$$
\mathbf{R} = \langle \mathbf{r} \mathbf{n} \rangle = \langle \mathbf{m} \mathbf{m} \rangle - \langle \mathbf{m} \mathbf{y} \rangle
$$
\n
$$
\mathbf{R} = \langle \mathbf{r} \mathbf{n} \rangle = \langle \mathbf{m} \mathbf{m} \rangle - \langle \mathbf{m} \mathbf{m} \rangle = \langle \mathbf{m} \mathbf{m} \rangle
$$
\n
$$
\mathbf{R} = \langle \mathbf{r} \mathbf{n} \rangle = \langle \mathbf{m} \mathbf{m} \rangle
$$

Adjust process noise **Q** to match the solution of the "Lyapunov Equation" to that of the <**nn** *^T*> estimate.

1 $\mathbf{P} \mathbf{a}$ as $\mathbf{t} \rightarrow \infty$, $\mathbf{P}_t = \mathbf{A} \mathbf{P}_{t-1} \mathbf{A}^T + \mathbf{G} \mathbf{Q} \mathbf{G}^T$ Then choose **Q** such that **HPH**^T \approx \langle **nn**^T

I.Fukumori/JPL

What are we really solving? How can we solve it?

OUTLINE

- 1. What is assimilation really solving?
	- a) Filtering vs Smoothing
	- b) Observability & Controllability
	- c) Data Error and Model Error
- 2. Examples and Practical Issues; e.g., ECCO
	- a) A hierarchical approach
	- b) Practical derivation and implementation
	- c) Assessment

"Estimating the Circulation and Climate of the Ocean"

http://www.ecco-group.org

GOAL: Advance data assimilation into an operational tool to help understand ocean circulation.

- 3 groups: JPL, Scripps Institution of Oceanography, Massachusetts Institute of Technology.
- Employ as much observations as possible.
- State-of-the-art ocean general circulation model.
- Advanced assimilation (Kalman filter/smoother, adjoint, Green's function) characterized by its consistency.

I.Fukumori/JPL

ECCO Near Real-Time Analysis

- **IPLODAP (External) Netscape** Edit View Go Communicator Help \mathbf{z} $\frac{\Delta t}{Home}$ Search Netscape $\overline{\mathbf{E}}$ \mathbf{f} Print Security Relond C Bookmarks & Location **Pro //12814933210/ovto** $+$ View the NASA Jet Propulsion Laboratory Portal **JPL Ocean Data Assimilation** A Member of ECCO Consortium (JPL | MIT | SIO) Home | Help •Ocean State Estimate for April 19, 2004 2004/04/19 s sea level ^c temperature c salinity c east-west velocity o north-south velocity depth $\boxed{0}$ = (m) view Sea level anomaly (cm) relative to average seasonal cycle **Project Description Access to Products Publications** Meeting Posters Distributed Oceanographic Data System (DODS ECCO - JPL Ocean Data Assimilation Project - Microsoft Internet Explor $\frac{1}{2}$ File Edit View Favorites Tools Help +Back ▼ → → ◎ ② △ | ◎Search ④Favorites ④Media ③ | 23 → ③ 回 → 国 ◎ 画 Address (a) http://ecco.jpl.nasa.gov/las/servlets/dataset $\overline{\cdot}$ \approx Go ECCO - JPL Ocean Data Assimilation Project $\boxed{60}$ Search $2.3.3.4.2.1$ noon
pare
nun Click on a dataset to continue or an \bigoplus for information about a dataset Datasets **Select dataset:** Variables **O** Assimilation (Adjoint) 1997-2000 **O** Assimilation (Kalman Filter) 1993 present Constraints **D** Assimilation (Smoothed Wind-Driven) 1993-2004 Output **O** Simulation 1980-pres Output Options Previous Outpu Define variable Abou LAS III Version 63 Show variables
	- Nowcasts every 10-days.
	- Highlights @
		- http://ecco.jpl.nasa.gov/external
		- SSH, Temperature profiles & timemean sea level.
		- Controls: winds, heat & fresh water fluxes and mixing parameters.
			- Green's function, Kalman filter & RTS Smoother, Adjoint method)
			- • LAS server from 1993 to present: http://www.eccogroup.org/las

ECCO Routine Assimilation

MIT General Circulation Model (MITgcm)

(Marshall et al., 1997, *JGR*)

- Nonlinear primitive equation model
- Adaptable to massively parallel supercomputers
- Advanced physics (e.g., KPP & GM mixing)
- Global, high resolution
	- $\sqrt{80}$ °S~ 80°N
	- \checkmark 0.3° × 10m in tropics

I.Fukumori/JPL

Approximate Kalman filter & Smoother

Kalman filter estimate :

$$
\hat{\mathbf{x}}_t^f = \mathbf{A}\hat{\mathbf{x}}_{t-1}^a + \mathbf{G}\mathbf{u}_{t-1}
$$
\n
$$
\hat{\mathbf{x}}_t^a = \hat{\mathbf{x}}_t^f + \mathbf{P}_t^f \mathbf{H}^T \left(\mathbf{H} \mathbf{P}_t^f \mathbf{H}^T + \mathbf{R}_t \right)^{-1} \left(\mathbf{y}_t - \mathbf{H} \hat{\mathbf{x}}_t^f \right)
$$
\n
$$
\mathbf{P}_t^f = \mathbf{A} \hat{\mathbf{P}}_{t-1}^a \mathbf{A}^T + \mathbf{G} \mathbf{Q} \mathbf{G}^T
$$
\n
$$
\hat{\mathbf{P}}_t^a = \mathbf{P}_t^f - \mathbf{P}_t^f \mathbf{H}^T \left(\mathbf{H} \mathbf{P}_t^f \mathbf{H}^T + \mathbf{R}_t \right)^{-1} \mathbf{H} \mathbf{P}_t^f
$$

 t *t* $\left\{ \frac{1}{t} \right\}$ *t* $\left\{ \frac{1}{t} \right\}$ *t* $\left\{ \frac{1}{t} \right\}$

The explicit evaluation of **P** ("Riccati Equation") is impractical because of their large dimension and time-evolution.

ECCO employs three approximations of **P**;

- 1. Time-asymptotic
- 2. Reduced-state
- 3. Partitioned

What are we really solving? How can we solve it?

OUTLINE

- 1. What is assimilation really solving?
	- a) Filtering vs Smoothing
	- b) Observability & Controllability
	- c) Data Error and Model Error
- 2. Examples and Practical Issues; e.g., ECCO
	- a) A hierarchical approach
	- b) Practical derivation and implementation
	- c) Assessment

1. Errors due to <u>wind error</u> (process noise) are estimated.

2. Partition global domain into <u>8 different regions</u> (basins)

- \bullet Global barotropic cell
- \bullet 7 Regional baroclinic cells

- I. Tropical Indian (>40 ºS)
- II.Tropical Pacific (40 \textdegree S ~ 40 \textdegree N)
- III. Tropical Atlantic (40 ºS ~ 40 ºN)
- IV. South Pacific (<20 ºS)
- V. South Atlantic-Indian (<20 ºS)
- VI. North Pacific (>20 ºN)
- VII. North Atlantic (>20 ºN)

3. State and control reduction: Horizontal velocity and vertical displacement in terms of vertical dynamic modes on a coarse horizontal grid. The reduced control is wind error on the same coarse grid.

3b. Horizontal Interpolation: **O**

"Objective Mapping" (Bretherton et al., 1976, *DSR*)

Given m<u>easurement</u>s φ , the underlying field $θ$ is estimated by

$$
\frac{\theta = CA^{-1}\varphi}{\text{is is just another}} \quad \text{where} \quad C = \left\langle \theta \varphi^T \right\rangle \quad A = \left\langle \varphi \varphi^T \right\rangle
$$
\n
$$
\text{least-squares} \quad \frac{\hat{\mathbf{a}} = \mathbf{a}_0 + \mathbf{R}_{\mathbf{aa}} \mathbf{E}^T \left(\mathbf{ER}_{\mathbf{aa}} \mathbf{E}^T + \mathbf{R}_{\mathbf{bb}} \right)^{-1} \left(\mathbf{b} - \mathbf{E} \mathbf{a}_0 \right)}{\mathbf{a}_0 + \mathbf{R}_{\mathbf{aa}} \mathbf{E}^T \left(\mathbf{ER}_{\mathbf{aa}} \mathbf{E}^T + \mathbf{R}_{\mathbf{bb}} \right)^{-1} \left(\mathbf{b} - \mathbf{E} \mathbf{a}_0 \right)}
$$

Horizontal interpolation, **O**, is defined as objective mapping from a coarse grid system to the model grid system.

Example of mapping of a perturbation at one coarse grid point;

This

- Global barotropic cell: global 6^o× 6^o grid, barotropic UVH (966 [×]3=2898 variables)
- Regional baroclinic cells: regional 5º× 3º grid, baroclinic UVD (5 gravest modes)
	- I.. Tropical Indian (>40 ºS ; 308 \times 3 \times 5=4620)
	- II.. Tropical Pacific (40 ºS ~ 40 ºN; 787 $\times 3$ $\times 5$ =11805)
	- III.. Tropical Atlantic (40 ºS ~ 40 ºN; 350 $\times3$ $\times5$ =5250)
	- IV.South Pacific (<20 ºS ; 633 [×]3 [×]5=9495)
	- V.. South Atlantic-Indian (<20 $^{\sf o}{\sf S}$; 664 $\times3$ $\times5{=}9960)$
	- VI. North Pacific (>20 ºN; 271 [×]3 [×]5=4065)
	- VII.. North Atlantic (>20 ºN; 198 $\times3$ $\times5$ =2970)

In comparison, the model's state dimension is 8 million (2 million grid points).

4. Compute **P** by formulating and explicitly deriving partitioned reduced-state models.

4a. Formulate partitioned reduced-state models.

The reduced-state must form a closed system.

A state and control perturbation δ **x**, δ **u** can be written as $\delta x = Bx' + Nn$ Given the model $\mathbf{x}_{t+1} = \Im(\mathbf{x}_t, \mathbf{u}_t)$ the perturbation satisfies $\delta \mathbf{x}_{_{t+1}} = \mathfrak{I}\big(\tilde{\mathbf{x}} + \delta \mathbf{x}_{_{t}}\big)$ $\mathbf{x}_{_t}, \mathbf{u}_{_t})$ **x** $\tilde{\mathbf{x}} + \delta \mathbf{x}_{t}, \tilde{\mathbf{u}} + \delta \mathbf{u}_{t}$) – $\mathfrak{I}(\tilde{\mathbf{x}}, t)$ $-\Im(\tilde{\mathbf{x}},\tilde{\mathbf{u}})$ or $\mathbf{Bx}'_{t+1} + \mathbf{N}\mathbf{n}_{t+1} = \Im(\tilde{\mathbf{x}} + \mathbf{Bx}'_t + \mathbf{N}\mathbf{n}_t, \tilde{\mathbf{u}} + \hat{\mathbf{B}}\mathbf{u}'_t + \hat{\mathbf{N}}\mathbf{m}_t) - \Im(\tilde{\mathbf{x}}, \tilde{\mathbf{u}})$ $\Big($) $n_{\text{rel}} = \Im[X]$ $\tilde{\mathbf{x}} + \mathbf{B}\mathbf{x}' + \mathbf{N}\mathbf{n}_\perp, \tilde{\mathbf{u}} + \mathbf{B}\mathbf{u}' + \mathbf{N}\mathbf{m}_\perp) - \mathfrak{I}(\tilde{\mathbf{x}}, \tilde{\mathbf{u}})$: **B**, **N** : range & null space of reduced-state δ $\mathbf{u} = \hat{\mathbf{B}} \mathbf{u}' + \hat{\mathbf{N}} \mathbf{m}$: The control equivalent of state reduction.

$$
\mathbf{Bx}'_{t+1} + \mathbf{N}\mathbf{n}_{t+1} = \mathfrak{I}(\tilde{\mathbf{x}} + \mathbf{B}\mathbf{x}'_t + \mathbf{N}\mathbf{n}_t, \tilde{\mathbf{u}} + \hat{\mathbf{B}}\mathbf{u}'_t + \hat{\mathbf{N}}\mathbf{m}_t) - \mathfrak{I}(\tilde{\mathbf{x}}, \tilde{\mathbf{u}})
$$

Multiplying by the pseudo inverse of $\textbf{B}, \textbf{B}^* \ \ (\textbf{B}^* \textbf{B} = \textbf{I}, \ \textbf{B}^* \textbf{N} = 0)$ $\mathbf{B}^* \mathbf{N} = 0$

$$
\mathbf{x}'_{t+1} = \mathbf{B}^* \left(\mathfrak{I} \left(\tilde{\mathbf{x}} + \mathbf{B} \mathbf{x}'_t + \mathbf{N} \mathbf{n}_t, \tilde{\mathbf{u}} + \hat{\mathbf{B}} \mathbf{u}'_t + \hat{\mathbf{N}} \mathbf{m}_t \right) - \mathfrak{I} \left(\tilde{\mathbf{x}}, \tilde{\mathbf{u}} \right) \right)
$$

For this to be a model for the reduced-state we cannot have dependency on the null space, \mathbf{Nn}_t , $\mathbf{\hat{N}m}_t$;

$$
\mathbf{x}'_{t+1} = \mathbf{B}^* \left(\mathfrak{I} \left(\tilde{\mathbf{x}} + \mathbf{B} \mathbf{x}'_t, \tilde{\mathbf{u}} + \hat{\mathbf{B}} \mathbf{u}'_t \right) - \mathfrak{I} \left(\tilde{\mathbf{x}}, \tilde{\mathbf{u}} \right) \right)
$$

This would be achieved if null (range) space perturbations \mathbf{r} emain in the null (range) space, because $\mathbf{B}^*\mathbf{N}=0$; i.e., state reduction must form a closed-system for it to be effective.

4b. Derive partitioned reduced-state model in **explicit matrix form** to facilitate error evaluation.

$$
\mathbf{x}'_{t+1} = \mathbf{B}^* \left(\mathfrak{I} \left(\tilde{\mathbf{x}} + \mathbf{B} \mathbf{x}'_t, \tilde{\mathbf{u}} + \hat{\mathbf{B}} \mathbf{u}'_t \right) - \mathfrak{I} \left(\tilde{\mathbf{x}}, \tilde{\mathbf{u}} \right) \right)
$$

$$
\mathbf{x}'_{t+1} = \mathbf{A}' \mathbf{x}'_t + \mathbf{G}' \mathbf{u}'_t
$$

A' and G' can be derived as **coarse grain Green's functions**;

$$
(\mathbf{A}')_i = \mathbf{A}'\mathbf{e}_i + \mathbf{G}'\mathbf{0} = \mathbf{B}^* (\Im(\tilde{\mathbf{x}} + \mathbf{B}\mathbf{e}_i, \tilde{\mathbf{u}}) - \Im(\tilde{\mathbf{x}}, \tilde{\mathbf{u}}))
$$

$$
(\mathbf{G}')_i = \mathbf{A}'\mathbf{0} + \mathbf{G}'\mathbf{e}_i = \mathbf{B}^* (\Im(\tilde{\mathbf{x}}, \tilde{\mathbf{u}} + \hat{\mathbf{B}}\mathbf{e}_i) - \Im(\tilde{\mathbf{x}}, \tilde{\mathbf{u}}))
$$

Similarly, **H'** can be derived by,

$$
(\mathbf{H}')_{i} = \mathbf{H}'\mathbf{e}_{i} = H(\tilde{\mathbf{x}} + \mathbf{B}\mathbf{e}_{i}) - H(\tilde{\mathbf{x}})
$$

where *H* is a function defining model equivalent of data.

 $(\mathbf{A}')_i = \mathbf{A}' \mathbf{e}_i + \mathbf{G}' \mathbf{0} = \mathbf{B}^* (\mathfrak{I} (\tilde{\mathbf{x}} + \mathbf{B} \mathbf{e}_i, \tilde{\mathbf{u}}) - \mathfrak{I} (\tilde{\mathbf{x}}, \tilde{\mathbf{u}}))$ $\tilde{\mathbf{x}} + \mathbf{B}\mathbf{e}_{i}$, $\tilde{\mathbf{u}}$) $-\Im{\tilde{\mathbf{x}}}$, $\tilde{\mathbf{u}}$

Implementing B^{*} requires some considerations;

1. \bullet is sparse but \bullet^* is not. Thus, we do $\mathbf{O}^* = \left(\mathbf{O}^T\mathbf{O}\right)^{-1}\mathbf{O}^T$ **O O**

 $\left(\mathbf{O}^T\mathbf{O}\right)^{-1}$ is small and is, therefore, explicitly computed and stored. O^T is implemented as the transpose (adjoint) of **O**.

2. η is evaluated by time-integrating vertical velocity instead of inverting

$$
\delta \mathbf{T} = \frac{\partial \mathbf{T}}{\partial z} \delta \mathbf{\eta}, \quad \delta \mathbf{S} = \frac{\partial \mathbf{S}}{\partial z} \delta \mathbf{\eta}
$$

which is singular where stratification is weak.

4c. Derive **representative** time-asymptotic limit of **P**.

- Stationary solution to the Riccati Equation requires a stationary system, i.e., time-invariant **A'**, **G'**, **H'**, **Q**, **R**.
	- Approximate with a 3-day assimilation cycle;

$$
\mathbf{x}'_{t+1} = \mathbf{A}'\mathbf{x}'_t + \mathbf{G}'\mathbf{u}'_t
$$

If $\mathfrak{I}(\tilde{\mathbf{x}}, \tilde{\mathbf{u}})$ is a 1-day integration. Then $\mathbf{A}'\!=\!\mathbf{C}'^3$ where $(\mathbf{C}')_i = \mathbf{B}^* (\mathfrak{I} (\tilde{\mathbf{x}} + \mathbf{B} \mathbf{e}_i, \tilde{\mathbf{u}}) - \mathfrak{I} (\tilde{\mathbf{x}}, \tilde{\mathbf{u}}))$ $\tilde{\mathbf{x}} + \mathbf{B}\mathbf{e}_{i}$, $\tilde{\mathbf{u}}$) $-\Im{\tilde{\mathbf{x}}}$, $\tilde{\mathbf{u}}$

• Approximate all observing operation **H'** and **R** with representative observations; e.g.,

Assume a full 10-day cycle of altimeter observations is available every 3-days but with 3.3 times the error variance.

e.g., TOPEX/POSEIDON and Jason have a 3-day subcycle.

5. Implementation

- a) Actual assimilation is conducted every 6-hours using all observations taken within \pm 3-hours of the assimilation instant.
- b) Filtering is performed using an alternate formulation of the Kalman gain;

 $\left(\mathbf{H}\mathbf{P}_t^f\mathbf{H}^T + \mathbf{R}_t\right)^{-1}\left(\mathbf{y}_t - H\left(\mathbf{\hat{x}}_t^f\right)\right)$ $\Big($ ˆ f **H**^T (**HP**^f**H**^T + **R**) \int (**v** $-H$ ($\hat{\mathbf{x}}$ ^f t *t* $\left\{ \begin{array}{ccc} 1 & \mathbf{1} &$ *a* \mathbf{K} -filter increment $\Delta \hat{\mathbf{x}}_t^a = \mathbf{P}_t^f \mathbf{H}^T \left(\mathbf{H} \mathbf{P}_t^f \mathbf{H}^T + \mathbf{R}_t \right)^{-1} \left(\mathbf{y}_t - H \right) \hat{\mathbf{x}}_t^g$

filter gain

$$
\mathbf{P}_t^f \mathbf{H}^T \left(\mathbf{H} \mathbf{P}_t^f \mathbf{H}^T + \mathbf{R}_t \right)^{-1} = \mathbf{P}_t^a \mathbf{H}^T \mathbf{R}_t^{-1}
$$

This form minimizes computations involving **P** and simplifies the matrix inversion that is needed.

 $\Big($ $\mathbf{\hat{x}}_{t}^{a}=\mathbf{P}_{t}^{a}\mathbf{H}_{t}^{T}\mathbf{R}_{t}^{-1}\Big(\mathbf{y}_{t}-H\big(\mathbf{\hat{x}}_{t}^{f}\big)\Big)$ $\mathbf{H}^T \mathbf{R}^{-1} (\mathbf{v}_i - H(\hat{\mathbf{x}}^f))$ *t* \mathbf{t} \mathbf{t} \mathbf{t} \mathbf{t} \mathbf{t} \mathbf{t} \mathbf{t} $\mathbf{F}_t^a = \mathbf{P}_t^a \mathbf{H}_t^T \mathbf{R}_t^{-1} \left(\mathbf{y}_t - H \right) \mathbf{\hat{x}}_t^T$ $\Delta \hat{\mathbf{x}}_t^a = \mathbf{P}_t^a \mathbf{H}_t^T \mathbf{R}_t^{-1} (\mathbf{y}_t - H) \hat{\mathbf{x}}_t$ K-filter increment

c) The K-filter increment is sum of partitioned increments;

$$
\Delta \hat{\mathbf{x}}_t^a = \sum_i \mathbf{B}_i \Delta \hat{\mathbf{x}}_{t,i}^{\prime a}
$$
\nwhere $\Delta \hat{\mathbf{x}}_{t,i}^{\prime a} = \mathbf{P}_i^{\prime a} \mathbf{H}_{t,i}^{\prime T} \mathbf{R}_t^{-1} \left(\mathbf{y}_t - H \left(\hat{\mathbf{x}}_t^f \right) \right)$
\nOr, $\Delta \hat{\mathbf{x}}_{t,i}^{\prime a} = \mathbf{P}_i^{\prime a} \mathbf{B}^T \mathbf{H}_{t,i}^T \mathbf{R}_t^{-1} \left(\mathbf{y}_t - H \left(\hat{\mathbf{x}}_t^f \right) \right)$
\nusing the adjoint of H in place of \mathbf{H}^T .

These calculations are performed from right to left; There is no explicit computation of product $\left[\mathbf{P}_i^{\prime a} \mathbf{H}_i^{\prime} \right]^{T} \mathbf{R}_i^{-1}$ because it is computationally inefficient. , $i \left(\begin{array}{cc} 1 & \cdots & t \\ \end{array} \right)$ $\mathbf{P}'^a \mathbf{H}'$ \cdot \mathbf{R} $\overline{}$

- d) Anomalies are assimilated for sea level and temperature profiles because of
	- i.Unknown marine geoid,
	- ii.Temporally uncorrelated process noise assumption.

$$
\Delta \hat{\mathbf{x}}_t^{\prime a} = \mathbf{P}_i^{\prime a} \mathbf{H}_{t,i}^{\prime T} \mathbf{R}_t^{-1} \left(\left(\mathbf{y}_t - \overline{\overline{\mathbf{y}}} \right) - \left(H \left(\hat{\mathbf{x}}_t^f \right) - \overline{\overline{\mathbf{m}}} \right) \right)
$$

where \bar{y} and \bar{m} are time-means of data and model (simulation) equivalent of data mean, respectively.

This choice of reference amounts to assimilating temporal anomalies of the data without altering the model time-mean.

e) Integrate original model after state and diagnostic variables are updated; partitioned reduced-state model is never used to integrate state.

f) Smoothing is performed

i.. Derive smoothed wind correction by,

$$
\Delta \mathbf{\hat{u}}_i^s = \mathbf{Q}_i \mathbf{G}^{\mathrm{T}} \left(\mathbf{G} \mathbf{Q}_i \mathbf{G}^{\mathrm{T}} + \mathbf{A} \mathbf{P}_i^a \mathbf{A}^{\mathrm{T}} \right)^{-1} \left(\mathbf{\hat{x}}_{i+1}^s - \mathbf{A} \mathbf{\hat{x}}_i^a - \mathbf{G} \mathbf{u}_i \right)
$$

 $\mathbf{TP}^{f^{-1}}\Big(\Delta \mathbf{\hat{x}}_{i+1}^s + \Delta \mathbf{\hat{x}}_{i+1}^a \Big)$ ˆˆ $f^{-1}\left(\Delta \mathbf{\hat{x}}_{i+1}^s + \Delta \mathbf{\hat{x}}_i^a\right)$ − $+1$ l + $= \mathbf{Q}\mathbf{G}^{\mathrm{T}}\mathbf{P}^{\mathrm{T}}$ $\left(\Delta\hat{\mathbf{x}}_{i+1}^{s} + \Delta\hat{\mathbf{x}}\right)$

In partitioned reduced-state;

$$
\Delta \hat{\mathbf{u}}_t^s = \sum_{i} \hat{\mathbf{B}}_i \Delta \hat{\mathbf{u}}_{t,i}^s
$$

where
$$
\Delta \hat{\mathbf{u}}_{t,i}^s = \mathbf{Q}_i^{\prime} \mathbf{G}_i^{\prime T} \mathbf{P}_i^{\prime f^{-1}} \Big(\Delta \hat{\mathbf{x}}_{t+1,i}^{\prime s} + \Delta \hat{\mathbf{x}}_{t+1,i}^{\prime a} \Big)
$$

ii. Re-run model simulation with smoother corrected wind because, $\mathbf{\hat{x}}_{t+1}^a + \Delta \mathbf{\hat{x}}_{t+1}^s \neq \mathfrak{I}(\mathbf{\hat{x}})$ $\hat{\mathbf{x}}_{t+1}^a + \Delta \hat{\mathbf{x}}_{t+1}^s \neq \Im\left(\hat{\mathbf{x}}_t^a + \Delta \hat{\mathbf{x}}_t^s, \mathbf{u}_t + \Delta \hat{\mathbf{u}}_t^s\right)$

due to approximations, but

$$
\hat{\hat{\mathbf{x}}}_{t+1}^{s} = \mathfrak{I}\left(\hat{\hat{\mathbf{x}}}_{t}^{s}, \mathbf{u}_{t} + \Delta \hat{\mathbf{u}}_{t}^{s}\right)
$$

Skill & Test of Assimilation

Data not explained by model

Assimilation should make models closer to data, both assimilated and independent.

Data variance explained by models; $\left\langle \mathbf{y}\mathbf{y}^T \right\rangle - \left\langle (\mathbf{y}-\mathbf{H}\mathbf{x})(\mathbf{y}-\mathbf{H}\mathbf{x})^T \right\rangle$

Skill & Test of Assimilation

Model-data residuals should be comparable to expectation

Skill & Test of Assimilation

Comparison to independent data: ECCO assimilation explains observed polar motion better than ocean simulation does.

Summary

- 1. Data assimilation is an inverse problem of observations & model with state & model controls as unknowns;
- 2. Data constraint error includes model representation error,

- 3. Kalman filters and RTS smoothers are recursive leastsquares inversions,
- 4. Kalman filters invert observations; RTS smoothers invert model,

Summary

- 5. Smoother's control estimates are fundamental in solving the assimilation problem, the assimilation problem,
- 6. Temporal evolution of filters is not physically consistent, but that of smoothers is,

7. Various approximations & simplifications are possible that make solving the assimilation problem tractable; ECCO KFS.

Future Challenge

Modeling different model error sources by identifying effective corresponding reduced-state approximations **B** is a central issue in data assimilation.

Controls need explicit modeling for estimation;

$$
\overline{\mathbf{x}}_{t+1} = A(\overline{\mathbf{x}}_t, \hat{\mathbf{u}}_t^0) + \left\{ \Pi L(\mathbf{w}_t, \mathbf{v}_t) - A(\Pi \mathbf{w}_t, \hat{\mathbf{u}}_t^0) \right\}
$$

= $A\overline{\mathbf{x}}_t + G\overline{\mathbf{u}}_t$ True evolution
True evolution
true model state

What are the most effective basis functions **B** that would represent effects of various model error sources ?

Effective = Small dimension but closed system.

$$
\mathbf{x}'_{t+1} = \mathbf{B}^* \Big(\mathfrak{I} \Big(\tilde{\mathbf{x}} + \mathbf{B} \mathbf{x}'_t + \mathbf{N} \mathbf{n}_t, \tilde{\mathbf{u}} + \hat{\mathbf{B}} \mathbf{u}'_t + \hat{\mathbf{N}} \mathbf{m}_t \Big) - \mathfrak{I} \Big(\tilde{\mathbf{x}}, \tilde{\mathbf{u}} \Big) \Big) \approx \mathbf{B}^* \Big(\mathfrak{I} \Big(\tilde{\mathbf{x}} + \mathbf{B} \mathbf{x}'_t, \tilde{\mathbf{u}} + \hat{\mathbf{B}} \mathbf{u}'_t \Big) - \mathfrak{I} \Big(\tilde{\mathbf{x}}, \tilde{\mathbf{u}} \Big) \Big)
$$